

pebbles; this is covered by a 7 m thick mostly shaly, silty and sandy interval, followed by a 2 m thick conglomeratic level, which resembles the Cocriamont conglomerate in its type section in the rue Cocriamont (stop 2.7). This conglomeratic level contains besides pebbles of siltstone and sandstone also well-rolled large cobbles (up to 55 cm in length) of limestone, with up to 57% of CaCO<sub>3</sub>, rich in macrofauna, indicating an eroded carbonate platform nearby. Forty samples were taken for chitinozoan studies and in five formations the organic microfossils could be extracted. As usual in most of the Condros Inlier, the chitinozoans are dark to opaque and only moderately preserved. They have a concentration between 0.1 and 14.5 chitinozoans per gram of rock and a diversity of 1 to 10 species per sample (Fig. 30). The indicative species are illustrated on the Plates 6-9.

The biozonation with chitinozoans could corroborate the age given with graptolites for the Vitruval-Bruyère Formation and the base of the Criptia Group. For the other three formations a detailed age can now be proposed for the first time (see below). The Fosses Formation was dated earlier with trilobites and brachiopods as pre-Hirnantian Ashgill by Sheehan (1987) and Lespérance & Sheehan (1987), mainly because no Hirnantian brachiopods were encountered. Hence an broad age of early to middle Ashgill age was inaccurately adopted, but a pre-Hirnantian late Ashgill age was also possible according to these macrofossils.

The most important result from the chitinozoans is that the Fosses and Génicot formations around the *Étang du Diable* and in the *Parc de Sart-Eustache* section are middle and late Ashgill in age and for the uppermost levels possibly Rhuddanian (early Llandovery). The samples in which the chitinozoans are frequent and diversified can be correlated with several biozones of Baltoscandia (Nolvak & Grahn, 1993; Nolvak, 1999), although not all of the Baltoscandian biozones are recognised here. The chitinozoans in the basal sample of the Fosses Formation can be correlated with the local biozone 10 (Fig. 9) in the Brabant Massif, present in the Harelbeke borehole (Samuelsson & Verniers, 2000), correlatable with the Baltoscandian *Tanuchitina bergstroemi* Biozone, dated as late Vormsi or early Pirgu, middle Ashgill. This age for the basal part of the Fosses Formation is the same age or younger than the top of the Madot Formation, the unit with a similar facies in the Brabant Massif (see stop 2.2). The chitinozoan assemblage from a higher level of the formation in the top of the quarry near the *Étang du Diable* can be correlated with the *Conochitina rugata* Biozone of Baltoscandia (Nolvak & Grahn, 1993; Nolvak, 1999), corresponding to the middle Pirgu (also middle Ashgill). Hence a broad middle part of the Ashgill is proposed for the whole Fosses Formation in its new def-

inition. Only a few but well preserved chitinozoans are present in the limestone cobbles of the Cocriamont conglomerate and indicate a late Ashgill age. The assemblage is very comparable with the assemblage in a shaly level about 4 m below that conglomerate. Samples from above the upper conglomerate contain a different chitinozoan assemblage possibly correlatable with the *Conochitina scabra* Biozone of Baltoscandia dated as late Porkuni (Nolvak & Grahn, 1993; Nolvak, 1999). The early Porkuni is considered to be the time of the Hirnantian glaciation, which on Gondwana is often expressed sedimentologically as two (very) coarse intervals. It is tempting to interpret the two coarse levels as the sedimentological expression of a sea level drop, resulting from the Hirnantian glaciation. It would be difficult to interpret otherwise the sudden presence of two coarse levels in between the dominantly fine clastic sedimentation in the Condros Inlier. Well preserved and rich chitinozoans in the base of the Criptia Group corroborate the late Rhuddanian age proposed with graptolites in this section. The transition between the Génicot Formation and the Criptia Group is however not observed here, due to an observation gap over 17 m in the section.

In the part of the Condros Inlier east of the city of Fosses-la-Ville, the area that we visit, the following formations are present and their descriptions are summarised from the recent revision by Verniers *et al.* (2001). The results of the study presented here however change already slightly the descriptions and ages of the formations in that revision. It appears that the Fosses Formation as defined now, e.g. on the new geological map of the area by Delcambre & Pingot (in press) corresponds only with what was previously included in the Bois de Presles Member, and that the recently defined Génicot Formation corresponds with what was called the Faulx-les-Tombes Member (Lassine *in* Martin, 1969a). The Cocriamont conglomerate would be situated in the middle part of the Génicot Formation. It can be mentioned here that after the excursion new samples of the Fosses and the Génicot Formations have been studied for chitinozoans in seven sections of the Puagne area (Vanmeirhaeghe & Verniers, submitted).

The new stratigraphical results presented above, can offer an explanation for the apparent contradiction in sedimentological interpretation for the deposition of the Fosses Formation (and the Génicot Formation) by Tourneur *et al.* (1993). Either a deposition as bioclastic turbidites with interbedded shales near a platform-ramp margin was suggested or a regressive event on a shelf followed by a transgressive sequence. Due to lack of a type section the exact position of samples in the lithostratigraphy and macrofossils was largely unknown. A



revision of the studied sedimentological samples and macrofossils should relocate these in the section. Because of the drastic environmental changes involved due to the glaciation-induced sea-level drops both sedimentological interpretations in Tourneur *et al.* (1993) are now possible, and not mutually exclusive any more.

### 7.2. Vitriual-Bruyère Formation

Description. Blackish, silky and fine micaceous shale with quartzite intercalations and graptolites (Maillieux, 1926); micaceous pelitic unit (Michot, 1928); the higher part contains clayey quartzitic sandstone in thick beds, covering black or blue shale with intercalations of black or dark coloured mudstone, siltstone and fine sandstone. The sedimentology has not been studied yet. The stratotype (stop 2.6) lies in the outcrops in the river bed of or along the Ruisseau Le Treko, west of the hamlet La Bruyère, 500 to 1050 m east of the church of the village Vitriual. Possibly the Oxhe Formation is a lateral equivalent or synonym for the unit in the Oxhe Inlier, situated 52 km to the NE.

Thickness: Difficult to estimate, except for the upper part with quartzite beds: 20-30 m.

Age. Graptolites in a locality at Vitriual-Bruyère in the upper part of the unit indicate the *Climacograptus peltifer* Biozone (Maillieux, 1933). A restudy of the fauna by Bulman (1950) confirmed the presence of the *Climacograptus peltifer* Biozone (now equivalent to the *Diplograptus foliaceus* Biozone), which is situated in the uppermost Costonian to lower Longvillian substages (uppermost Aurelucian and most of the Burrellian stages, Caradoc) (Fortey *et al.*, 1995). The chitinozoans in the present study indicate the same assemblage, and the sampled levels belong to the local biozone 2 (Fig. 30) with as age indicators *Cyathochitina calix* and *Laufeldochitina stentor*. Comparing the range of these species with Baltoscandia (Nolvak, 1999), a broad late Llanvirn to early Caradoc age is attributed. By the presence of the first species it is more or less correlatable with the local biozone 4 of the Brabant Massif present in the Rigenée Formation (Samuelsson & Verniers, 2000).

### 7.3. Basse-aux-Canes Formation

Description. Greenish grey or dark grey silty shale, often sandy siltstone, with irregular jointing. White mica often present. The siltstone weathers with dark or rusty patches of iron and manganese oxides, described as a characteristic shining blue colour on freshly broken surface. No sandstone beds observed. The formation overlays the Sart-Bernard Formation. The unit resembles slightly the Vitriual-Bruyère Formation, but has no sandstone beds; it is possibly a lateral facies change of

(a part of) that formation. The stratotype area lies in the southern part of the Puagne area, south-east of Sart-Eustache but a section is not defined yet (Michot, 1934, 1954; Martin, 1969a; Delcambre & Pingot, in press; Verniers *et al.*, 2001).

Thickness. Estimated 100 to 150 m.

Age. Without fossils it could only be estimated as Llanvirn or early Caradoc (Middle or Late Ordovician). Two samples with chitinozoans described in the present study form the local biozone 1 (Fig. 30). The range of *Belonechitina capitata* and *Conochitina dolosa* compared to Baltoscandia (Nolvak, 1999) narrow the age down from the late Llanvirn to middle Caradoc.

### 7.4. Fosses Formation

Description. From top to bottom green and dark green fine sandy shale, with trilobites and brachiopods; greywacke, calcareous shale with locally some limestone beds rich in brachiopods (Michot, 1934). Martin (1969a after Michot, 1927, 1934) divided the formation in two members: a lower Bois de Presles Member with clayey limestone, calcareous shale, fossiliferous with brachiopods, cystoids and trilobites, in an alternation of thin-bedded muddy limestone and calcareous shale, in and in the upper part of the formation calcareous shale with thin limestone layers. The brachiopods and trilobites were described by Sheehan (1987) and Lespérance & Sheehan (1987), the crinoids, bryozoans, echinoderm debris, cystoids, molluscs and algae were described by Tourneur *et al.* (1993) and the corals by Servais *et al.* (1997). The base is not yet observed. Several sections were informally designated as stratotype: the northern bank of the Fuelle and Rosière rivers one km east of the city of Fosses (Malaise, 1900) and the area between the city of Fosses and the village of Sart-Eustache. In the present study we propose the outcrops and the abandoned quarry north of the *Etang du Diable* and the section in the *Parc de Sart-Eustache* as the stratotype section (Fig. 31).

Thickness. About 50 m (Tourneur *et al.*, 1993); more than 38 m in the *Parc de Saint-Eustache* section; estimated at more than 75 m.

Age. Based on trilobites and brachiopods: pre-Hirnantian Ashgill (Sheehan, 1987; Lespérance & Sheehan, 1987); based on chitinozoans: local biozones 3 to part of 6 (Fig. 30); Baltoscandian biozones of Nolvak & Grahn (1993) and Nolvak (1999): at the base of the formation *Tanuchitina bergstroemi* Biozone (upper Vormsi to lower Pirgu) and in the middle of the formation the *Conochitina rugata* Biozone (middle Pirgu), middle to late Ashgill.

### 7.5. Génicot Formation

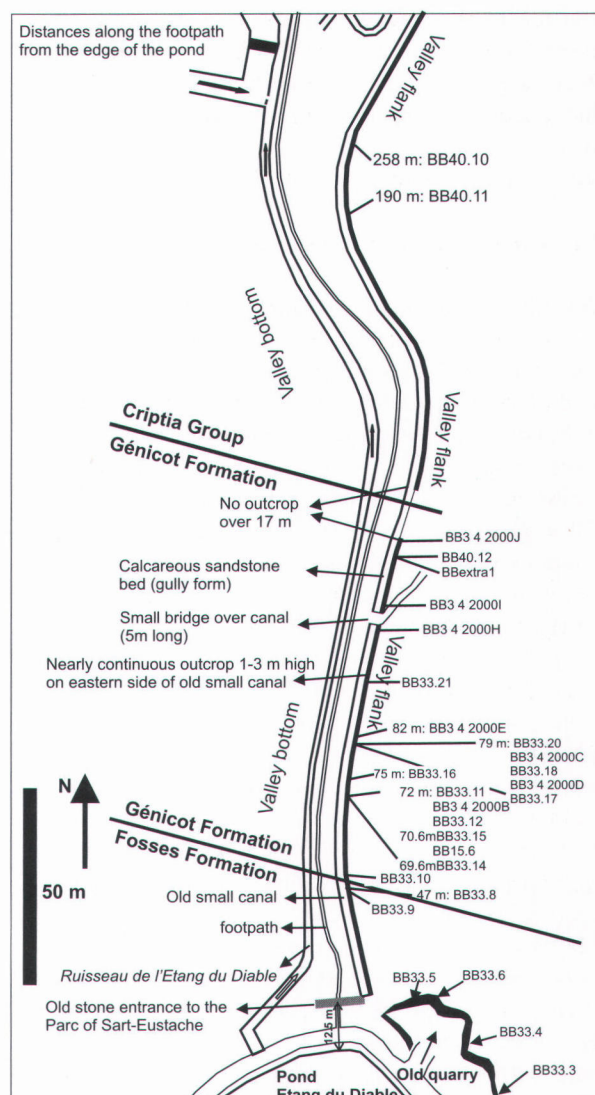
**Description.** Black shale, dark grey laminated siltstone and often thin beds of medium grey sandstone. In the middle part two sandstone beds occur, separated by a sandy-shaly unit from an overlying conglomeratic bed with large limestone cobbles. The latter bed was sometimes described as the Cocriamont conglomerate, a coarse crinoidal limestone above a basal conglomerate, which was considered by earlier authors (Michot, 1931) as the base of the Fosses Formation. High in the formation a few meters thick bed occurs with dark coloured microbioturbations, a few mm long (*Chondrites* sp.?) dispersed in the shale or siltstone. They were called “*schistes mouchetés*” by Lassine (unpublished in Martin, 1969a) or described as green sandy shale with on breaking surfaces blackish elliptic or fusiform spots and a rare macrofauna level (Martin, 1969a after Michot, 1927, 1934). The base of the formation is located above the highest rich macrofossil bearing horizon of the Fosses Formation. The top of the formation is formed by the highest siltstone/sandstone bed below the Criptia Group. It was considered by Martin (1969a) as the Faulx-les-Tombes Member, the upper member of the Fosses Formation. The *Parc de Sart-Eustache* section is proposed herein as the stratotype section.

**Thickness.** At least 85 m in the stratotype section (this study), but probably between 150 and 200 m (Delcambre & Pingot, in press).

**Age.** The chitinozoan assemblages allow the distinction between local biozones 6 (part) to 11 (Fig. 30). All of them contain the rare but distinct *Armoricochitina nigerica*, indicative for the late Ashgill. Above the conglomerate of Cocriamont level, the *Conochitina scabra* Biozone occurs, dated as late Porkuni by Nolvak & Grahn (1993) and Nolvak (1999), the part of the latest Ashgill after the Hirnantian glaciation. *Hercochitina* cf. *gamachiana* in this assemblage resembles the similar or same species in the local chitinozoan biozone 11 of the Brabant Massif in Samuelsson & Verniers (2000). If this can be confirmed by more studies it would place this biozone 11 also in the post-Hirnantian glaciation latest Ashgill. No chitinozoans were discovered yet in the top of the formation.

### 7.6. Criptia Group

**Description.** A group described as “Schistes de Criptia” by Michot (1928) and re-used by Delcambre & Pingot (in press) for a thick unit of shale and silty shale, homogeneous with various colours (greenish, grey, ochre) and a stratification difficult to distinguish. In the group a green to greenish grey soft shale unit can be differentiated. The fine-grained and compact sedimentation



**Figure 31.** Detailed location map of the samples for chitinozoan studies in the quarry near the *Etang du Diable* and in the *Parc de Sart-Eustache*, the proposed new type locality and section of the Fosses and the Génicot formations, changed after Billiaert (2000).

points to a deep shelf environment. The Génicot Formation is supposed to be the underlying unit. The covering unit is a more silty and darker coloured unnamed shale unit. The stratotype area is around the village of Sart-Eustache, but not defined yet.

**Thickness.** Unknown, tentatively estimated at several hundred meters.

**Age.** Only at the base well preserved and frequent chitinozoans are described in this study belonging to the *Belonechitina postrobusta* global Biozone of Verniers *et al.* (1995) of the Rhuddanian (early Llandovery).

Graptolites are found in higher levels but are still unstudied; an early Llandovery or possibly Wenlock to early Ludlow age is tentatively postulated (Delcambre & Pingot, in press; Verniers *et al.*, 2001).

## 8. Description of excursion stops DAY 2 afternoon

### Stop 2.6. *Vitrival-Bruyère section*

Location. Outcrops along a track and in the river bed of the Ruisseau Le Treko, west of the hamlet La Bruyère, 500 to 1050 m east of the church of the village Vitrival.

General structure. Subvertical bedding with stratification only visible in the decimetre to metre thick sandstone beds.

Lithostratigraphy. Vitrival-Bruyère Formation.

Lithology. Micaceous pelitic unit, with dark grey, black or blue, silty and fine micaceous shale with some thick quartzite intercalations and other beds of black or dark coloured mudstone, siltstone and fine sandstone. The sedimentology is unstudied yet.

Biostratigraphy. See above in chapter 7.2.

Remarks. The outcrop shows the difficulty to describe complete lithostratigraphical sections in the Condroz Inlier. The complexity of the Variscan deformation is enhanced by the suspected Quaternary mass movements along the steep hill side. The thickness of the formation cannot be determined. The lithology is however considered typical for the formation and chitinozoans allowed a corroboration of the earlier dating (see above in 7.2).

### Stop 2.7. *Cocriamont conglomerate type section*

Location. Rue de Cocriamont, Le Roux, Fosses-la-Ville. low escarpment in the west side of the road.

General structure. Steeply dipping beds, normal to overturned.

Lithostratigraphy. Cocriamont conglomerate level.

Lithology & sedimentology. See in 7.5.

Biostratigraphy. See in 7.4 and 7.5. One chitinozoan sample at 1.0 to 1.2 m above the conglomerate contains some specimens of species of the *Conochitina scabra* Biozone, suggesting tentatively the latest Ashgill.

### Stop 2.8. *Outcrop east of the "Etang du Diable", Sart-Eustache*

Location. Escarpment in an outcrop at the corner of the road, east of the "Etang du Diable".

General structure. Steeply dipping beds and northwards younging.

Lithostratigraphy. Fosses Formation (lower part).

Lithology. Typical (slightly) calcareous mudstone with macrofossils dispersed or in centimetre thick layers.

Sedimentology. Not studied yet here.

Biostratigraphy. The chitinozoans are rather rich and belong to the *Tanuchitina bergstroemi* Biozone, middle Ashgill (see in 7.5).

### Stop 2.9. *Abandoned quarry north the "Etang du Diable" and section in the "Parc of Sart-Eustache", Sart-Eustache*

Location. A continuous section can be followed from an abandoned quarry, along a public path, north of the "Etang du Diable" and in the private property of the Park of Sart-Eustache in a 750 m long nearly continuous section along a small disused canal (written permission should be asked to enter the property).

General structure. Steeply bedded normal to overturned beds with a main strike of N110-290. The sequence youngs to the north, evidenced by sedimentological structures.

Lithostratigraphy. In the first 250 m successively the Fosses and Gécicot formations, with the Cocriamont conglomerate and the Criptia Group is outcropping (Fig. 31).

Lithology & sedimentology. See above in 7.4, 7.5 and 7.6.

Biostratigraphy. See above; the chitinozoan succession indicates the presence of middle and late Ashgill to Rhuddanian.

Remarks. In this section the presence of the Hirnantian glaciation could be postulated for the first time in Belgium.

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